

Research Article

Spatial Patterns and Temporal Variability of Drought in Beijing-Tianjin-Hebei Metropolitan Areas in China

Wanyuan Cai,¹ Yuhu Zhang,¹ Qiuhua Chen,² and Yunjun Yao³

¹College of Resources Environment & Tourism, Capital Normal University, Beijing 100048, China

²School of Mathematical Sciences, Capital Normal University, Beijing 100048, China

³State Key Laboratory of Remote Sensing Science, College of Global Change and Earth System Science, Beijing Normal University, Beijing 100875, China

Correspondence should be addressed to Yuhu Zhang; zhang_yuhu@163.com

Received 4 July 2014; Revised 29 October 2014; Accepted 3 November 2014

Academic Editor: Francisco J. Tapiador

Copyright © 2015 Wanyuan Cai et al. This is an open access article distributed under the Creative Commons Attribution License, which permits unrestricted use, distribution, and reproduction in any medium, provided the original work is properly cited.

Drought identification and assessment are essential for regional water resources management. In this paper, the spatiotemporal characteristics of drought were evaluated based on monthly precipitation data from 33 synoptic stations during the period of 1960–2010. The percent of normal precipitation was applied to illustrate the driest years in Beijing-Tianjin-Hebei metropolitan areas (BTHMA) (1965, 1997, and 2002). The modified Reconnaissance Drought Index (RDI) was applied to capture the drought patterns and to estimate the drought severity at 33 meteorological stations. Agglomerative hierarchical cluster analysis (AHCA) and principal component analysis (PCA) were used to identify three different drought subregions R1, R2, and R3 based on the monthly precipitation values in BTHMA, which is located in southeast, north, and south of BTHMA, respectively. The year 1965 was the driest and 1964 was the wettest during the observed period. The characteristics of drought were analyzed in terms of the temporal evolution of the RDI-12 values and the frequency of drought for the three identified regions. The percentage of years characterized by drought was 13.73% for R1, 16.50% for R2, and 15.53% for R3. 66.91% of drought belongs to the near normal drought category. The obtained results can aid to improve water resources management in the area.

1. Introduction

In recent years, climate change and the growing global warming trend have aroused people's concern, which frequently caused the extreme events [1–3]. As one of the most serious disasters in the large of extreme weather events, drought has devastating impacts on water resources, the environment, and the human health in some regions, even all over the world [4, 5]. Drought is not only a complex natural hazard but a disaster [6], which is defined by the lack of precipitation [7]. Besides, regional drought has become one of the vital researches on regional studies of global change [8]. And it becomes important to study the drought distribution characteristics on the time and space of a region and what caused the drought [9, 10].

Drought is often represented in terms of drought variables [11], which include drought intensity, drought frequency, and duration. A large number of drought indices with various

complexities have been used in many areas all over the world for various purposes. Some of the most popular indices used in the past include the Palmer Drought Severity Index (PDSI) [12], the Rainfall Anomaly Index (RAI), the Soil Moisture Drought Index (SMDI), the Standardized Precipitation Index (SPI) [13], the Deciles, the Percent of Normal, the Crop Moisture Index (CMI), the Palmer Hydrological Drought Index (PHDI), the Surface Water Supply Index (SWSI) [11, 14], the Standardized Anomaly Index (SAI), and indices based on the Normalized Difference Vegetation Index (NDVI) [15]. Heim [16] summarized a comprehensive review of drought indices used in the United States for 20th century.

During the first decade of the 21st century, the Standardized Precipitation Index (SPI) was widely utilized, which involves only precipitation data and can be detected nearly everywhere. Recently a new index, the Reconnaissance Drought Index (RDI), was advanced [17, 18].

The percent of normal precipitation was applied to illustrate the driest years and the modified RDI was used to capture the drought patterns and estimate the drought severity. RDI is widely used and is gaining ground, mainly owing to its wickered data conditions and its high sensitivity and elasticity [19–23]. And it is based both on accumulative precipitation (P) and on potential evapotranspiration (PET). In addition, previous studies have detected that the use of different PET methods has no significant influence on RDI. This also supports the perspective that RDI is a vigorous drought index, not dependent on the PET calculation methods, which simplifies the process of calculation [15].

Numerous studies worldwide have been conducted to analyze the spatial and temporal characteristics of drought. Rossi et al. [24] focused on spatial aspects of drought by examining all drought characteristics. For this goal, the study quantified droughts at different locations using several types of hydrologic data from all the observation sites within the study area. To obtain the better understanding of spatial patterns, a predicting model was developed and applied. Clausen and Pearson [25] analyzed the relationship between duration and severity of the largest annual droughts at various locations by applying linear regression analysis. Moreover, the regional drought frequency analysis was performed to achieve more reliable results for study areas with limited or inadequate data available. Estrela et al. [26] studied the impact and the frequency of drought, as well as its pressures on water resources. They highlighted that precipitation across Europe has been reducing during the last three decades of the 20th century. As a result, the number of extreme dry periods was increased over the last decade of the 20th century. Besides, a lot of researches have been conducted to better estimate spatial patterns on drought intensity and duration. Yoo and Kim [27] investigated the vulnerability of an environment to drought based on soil moisture. The spatial-temporal patterns of drought were characterized by applying the empirical orthogonal function (EOF), which enables us to identify major styles of spatial variability. Gocic and Trajkovic [7] analyzed the evaluation of spatiotemporal characteristics of drought based on monthly precipitation data from meteorological stations. The percent of normal precipitation was applied to illustrate the driest years in Serbia and the Standardized Precipitation Index (SPI) and principal component analysis (PCA) were used to capture the drought patterns. Cluster analysis was applied to identify different drought subregions. The characteristics of drought were analyzed in terms of the temporal evolution of the SPI-12 values and the frequency of drought for the identified subregions.

Furthermore, there are some studies on evaluating the characteristics of drought for different periods and sites in China. For instance, Yuan and Wu [28] introduced the agricultural drought index (CSDI) and analyzed space and temporal changes of agricultural drought in the study area. By analyzing the CSDI values of 18 representative stations distributed in BTHMA from during the period of 1961–1990, four types of agricultural drought in this area were identified. Risk analysis on agricultural drought further showed the possibility of drought afflicted in agricultural production in

the area. Yan et al. [29] applied standardized precipitation index (SPI) as drought index and used precipitation from meteorological stations of China in the years of 1958 to 2007 to calculate the indices in each of seasons. Through applying Kriging interpolation to SPI values for each station all the values could be spatially and temporally comparable. Based on raster data of seasonal SPI, drought rate and drought probability were computed to demonstrate the spatial and temporal distribution characteristics of drought in Hebei province from the years of 1958 to 2007. Liu et al. [30] studied the spatial anomaly and temporal evolution characteristics of annual standard drought index by using EOF, the rotated empirical orthogonal function (REOF), wavelet analysis, and Mann-Kendall test based on the data of monthly precipitation and monthly average temperature of 589 meteorological stations over China from 1961 to 2009. The results showed that abnormality of the annual standard drought index over China was significant in ten areas. Among them, the climate became significantly dryer in seven regions and it became significantly wetter in 3 regions. There existed multiple time scale features over China for arid-wet change according to wavelet energy spectrum. However, a comprehensive analysis considering both precipitation and evapotranspiration in precipitation series and drought in regions as presented here is still lacking.

In this study, two drought indices—the percentage of normal precipitation and the Reconnaissance Drought Index (RDI)—were used. The percentage of normal precipitation was preliminarily applied to illustrate the driest years in Beijing-Tianjin-Hebei metropolitan areas, which is an effective index when applied to a single region [31]. The Reconnaissance Drought Index (RDI) was used as an input for a principal component analysis (PCA) to identify the drought patterns. It is reliable, since it calculates the aggregated difference between precipitation and the evapotranspiration. It is available under “climate instability” conditions, for checking the significance of various alterations of climatic factors related to water scarcity. For better reflecting the spatial heterogeneity of regional drought, the Reconnaissance Drought Index (RDI) was modified based on the climatic characteristics of the study area. In China, neither has the Reconnaissance Drought Index (RDI) been used in former researches nor have the cluster analysis and principal component analysis (PCA). Thus, the evaluation results based on modified RDI and PCA will provide some scientific assistance for decision makers when devising drought and water resources management policies to mitigate the adverse influence of drought.

The main objective of this study was (1) to calculate the percent of normal precipitation to illustrate the driest years in Beijing-Tianjin-Hebei metropolitan areas; (2) to estimate drought severity using RDI at the 12-month timescales at 33 meteorological stations; (3) to consider spatial and temporal variability of drought in Beijing-Tianjin-Hebei metropolitan areas during the period 1960–2010; (4) to identify subregions using PCA and cluster analysis and accomplish the characterization of the drought in the identified subregions.

2. Materials and Methods

2.1. Study Area and Data. Beijing-Tianjin-Hebei metropolitan areas (BTHMA), China's northernmost metropolitan region with its major cities, Beijing and Tianjin, are located in Hebei province and stretch from the municipalities of Beijing and Tianjin to the Bohai Sea. The study area comprises roughly 185,000 km². Southeast of BTHMA is mainly flat, while its northern areas consist of plateaus and mountains. The largest part of the area has the continental precipitation regime. The geographic features affecting the climate of BTHMA are the Inner Mongolia Plateau, the Hubei Plains, and the Bohai Sea. On average the maximum precipitation occurs in July and August and the minimum in February.

Monthly precipitation sets from 33 meteorological stations which have continuous record were acquired from meteorological data sharing service system of China (<http://cdc.cma.gov.cn>). The investigated time series were selected according to the availability and reliability of the data sets. Thus, a record length of 51 years (1960–2010) was considered, which is the maximum time period of precipitation data recorded covering all the 33 synoptic stations. The main information about the stations is presented in Table 1 and the geographical set of the stations is shown in Figure 1.

2.2. Methods

2.2.1. Percent of Normal Precipitation. Percent of normal precipitation is estimated by dividing actual precipitation by normal precipitation (at least 30-year mean period) and multiplying by 100%. Normal precipitation for a specific location is considered to be 100%, while the value of the index less than 100% means that there are drought conditions. As the simplest measure of precipitation, it is not useful for making decisions when used alone [31].

2.2.2. Modified Reconnaissance Drought Index (RDI). Tsakiris and Vangelis [32, 33] proposed the reconnaissance drought index (RDI), utilizing the ratios of precipitation over PET for different time scales, for representation over the region of interest. For the annual time scale RDI index is derived by first calculating α_0^i :

$$\alpha_0^i = \frac{\sum_{j=1}^{12} P_{ij}}{\sum_{j=1}^{12} PET_{ij}}, \quad i = 1(1)n, j = 1(1)12, \quad (1)$$

where P_{ij} and PET_{ij} are precipitation and potential evapotranspiration of the j th month of the i th year, respectively, and n is number of years for investigated data set. Generally, the Penman-Monteith equation is used to calculate PET; however, if required parameters are not available, it is recommended to use the Hargreaves-Samani equation [34]. Suitability of the Hargreaves-Samani equation has also been recommended in recent research works, for example, [15, 35]. Consequently, in the present research with limited data (only temperature), the Hargreaves-Samani equation [36] was applied to calculate PET. Considering that the use

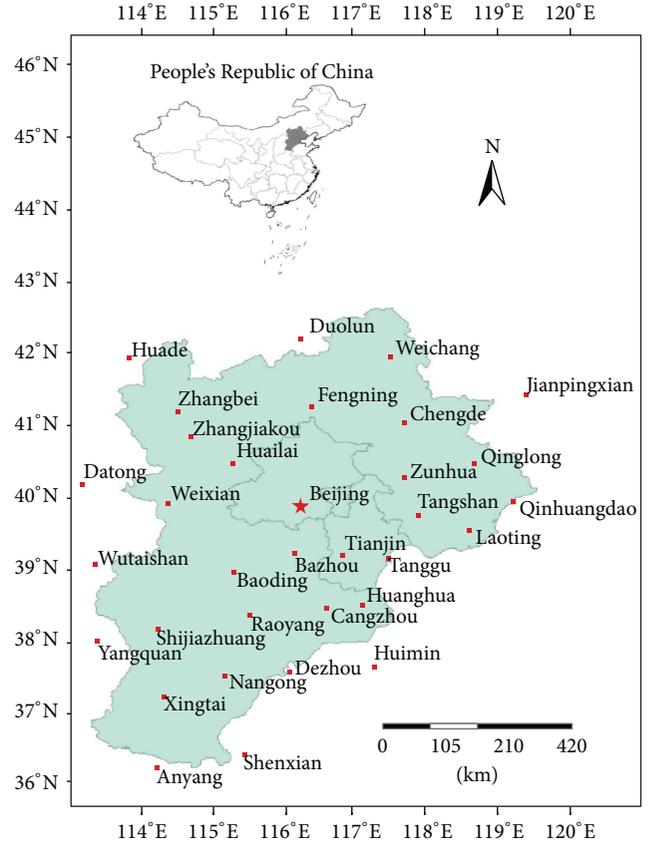


FIGURE 1: Spatial distribution of the 33 meteorological stations in Beijing-Tianjin-Hebei metropolitan areas map.

of different PET methods has no significant influence on RDI, we select the simplest Thornthwaite method with minimum data requirements [37]. However, the temperature is corrected using the effective temperature instead of the mean. The effective temperature is defined as $T_{ef} = 0.36(3T_{max} - T_{min})$ [38, 39].

Normalized RDI (RDI_n) is calculated as

$$RDI_n^{(i)} = \frac{\alpha_0^{(i)}}{\bar{\alpha}_0} - 1, \quad (2)$$

where $\bar{\alpha}_0$ is the arithmetic mean of α_0 values, computed for the n years of data.

Then, the standardized RDI (RDI_{st}) is computed in the way similar to SPI:

$$RDI_{st}^{(i)} = \frac{y_k^{(i)} - \bar{y}_k}{\hat{\sigma}_{y_k}}, \quad (3)$$

where $y_k^{(i)}$ is the $\ln(\alpha_0^{(i)})$, \bar{y}_k is the arithmetic mean, and $\hat{\sigma}_{y_k}$ is the standard deviation of y_k , assuming that the log-normal distribution is appropriate for α_0 [18]. Finally, the log-normal probabilities are transformed into Z normal values [40]. The standardized RDI behaves similar to the SPI and so is the interpretation of results. Therefore, the RDI_{st} can be compared to the same thresholds as the SPI.

TABLE 1: Geographical descriptions mean and standard deviation of annual precipitation time series of the synoptic stations used in the study.

Station name	Longitude (E)	Latitude (N)	Elevation (m)	Mean (mm)	Standard deviation (mm)
Beijing	116.47	39.80	31.30	536.32	172.07
Zhangbei	114.70	41.15	1393.30	386.94	68.47
Weixian	114.57	39.83	909.50	398.88	88.09
Shijiazhuang	114.42	38.03	81.00	525.89	172.28
Xingtai	114.50	37.07	77.30	515.19	176.86
Fengning	116.63	41.22	661.20	458.04	89.55
Weichang	117.75	41.93	842.80	433.49	88.21
Zhangjiakou	114.88	40.78	724.20	399.00	91.85
Huailai	115.50	40.40	536.80	378.98	80.58
Chengde	117.95	40.98	385.90	518.28	108.71
Zunhua	117.95	40.20	54.90	711.88	201.17
Qinglong	118.95	40.40	227.50	691.79	190.63
Qinhuangdao	119.52	39.85	2.40	634.01	179.22
Bazhou	116.38	39.12	9.00	511.18	184.55
Tangshan	118.15	39.67	27.80	605.37	160.43
Laoting	118.88	39.43	10.50	600.22	176.73
Baoding	115.52	38.85	17.20	517.57	195.07
Raoyang	115.73	38.23	19.00	519.99	153.49
Cangzhou	116.83	38.33	9.60	610.09	194.76
Huanghua	117.35	38.37	6.60	589.70	196.37
Nangong	115.38	37.37	27.40	478.54	142.07
Anyang	114.40	36.05	62.90	561.60	169.20
Jianpingxian	119.70	41.38	422.00	459.86	107.01
Huade	114.00	41.90	1482.70	307.79	66.02
Duolun	116.47	42.18	1245.40	375.24	70.05
Dezhou	116.32	37.43	21.20	567.40	182.87
Huimin	117.53	37.48	11.70	572.64	172.16
Shenxian	115.67	36.23	37.80	539.63	158.65
Datong	113.33	40.10	1067.20	370.24	84.31
Wutaishan	113.52	38.95	2208.30	748.29	183.55
Yangquan	113.55	37.85	741.90	541.07	148.24
Tianjin	117.07	39.08	2.50	536.45	147.41
Tanggu	117.72	39.05	4.80	575.59	183.33

The choice of the lognormal distribution is not constraining but it assists in devising a unique procedure instead of various procedures depending on the probability distribution function, which best fits the data. However, the hypothesis that the data of the RDI_n follow a lognormal distribution seems to be the most appropriate. In all examples analyzed during the establishment of the RDI, the goodness-of-fit tests confirmed that the lognormal distribution fits the data satisfactorily.

It should be emphasized that the RDI is based both on precipitation and on potential evapotranspiration. The mean initial index (α_0) represents the normal climatic conditions of the area and is equal to the Aridity Index as was proposed by the FAO.

Among others, some of the advantages of the RDI are as follows.

It is physically sound, since it calculates the aggregated deficit between precipitation and the evaporative demand

of the atmosphere. It can be calculated for any period of time. The calculation always leads to a meaningful figure. It can be effectively associated with agricultural drought. It is directly linked to the climatic conditions of the region, since for the yearly value it can be compared with the FAO Aridity Index. It can be used under "climate instability" conditions, for examining the significance of various changes of climatic factors related to water scarcity.

With advantages given above, it can be concluded that the RDI is an ideal index for the reconnaissance assessment of drought severity for general use giving comparable results within a large geographical area, such as the BTHMA.

It should be mentioned that usually droughts in the BTHMA are accompanied by high temperatures, which lead to higher evapotranspiration rates. Evidence for this has been produced from simultaneous monthly data of precipitation and evapotranspiration in BTHMA. From the cases analyzed it seems that about 90% of them comply with the previous

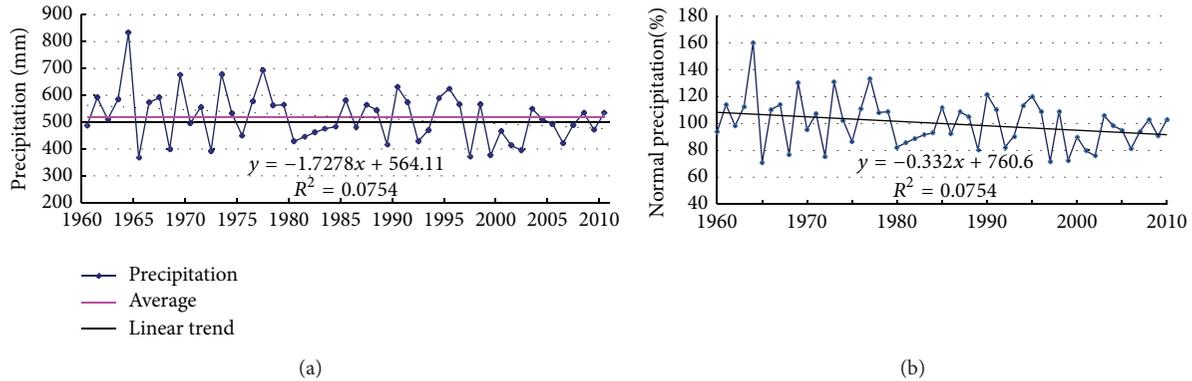


FIGURE 2: (a) Time series of total annual precipitation averaged over whole Beijing-Tianjin-Hebei metropolitan areas and (b) percent of normal precipitation index computed with respect to 1961–1990 climate normal.

TABLE 2: Drought classification of RDI.

Drought class	RDI value
Extremely wet	$RDI \geq 2.0$
Very wet	$1.5 \leq RDI < 2.0$
Moderately wet	$1.0 \leq RDI < 1.5$
Near normal	$-1.0 \leq RDI < 1.0$
Moderate drought	$-1.5 \leq RDI < -1.0$
Severe drought	$-2.0 \leq RDI < -1.5$
Extreme drought	$RDI < -2.0$

statement [41]. Therefore, the RDI is expected to be a more sensitive index than those related only to precipitation, such as the SPI.

According to Tigkas et al. [42], we divided the RDI into moderate, severe, and extreme classes for both dry and wet RDI as shown in Table 2. In this study, the 12-month timescales were used to monitor hydrological conditions and the impact of drought on the available water resources.

2.2.3. Principal Component and Cluster Analysis. Combination of techniques such as principal component analysis (PCA) [43] and cluster analysis (CA) [44, 45] can be applied to climate or drought regionalization [46–51]. PCA is a multivariate technique that reduces the dimensionality in a dataset and computes a set of new orthogonal variables with the decreasing order of importance named principal components (PCs) [52–54]. It is based on the estimation of the eigenvalues and eigenvectors from the characteristic equation. For this purpose, either the correlation or the covariance matrix of the observed variables can be used. Richman [55] defined six modes of PCA due to a different combination of time, objects, and attributes. In this study, the S-mode (data matrix with rows for the observations and columns for the stations) with the varimax orthogonal rotation method was applied to identify the spatial patterns of drought. The patterns defined in this way are named rotated loadings.

The rule of thumb [56] and scree plot [57] were applied to make the decision on how many principal components to retain for rotation. Bartlett's test of sphericity [58] and the Kaiser-Meyer-Olkin test [59] were performed to test the quality of the principal components. Then, the agglomerative hierarchical cluster analysis was applied on obtained rotated PC scores (RPC) using Ward's method [60] and Euclidean distance as the distance or similarity measure. According to Jolliffe [52], the CA is available when there is no clear group structure in the dataset. In this study, the CA is used to identify different drought subregions. This is in an agreement with the methodology applied by Raziei et al. [61], Santos et al. [62], and Martins et al. [51].

2.2.4. Inverse Distance Weighting. One of the most widely used deterministic methods in spatial interpolation is the inverse distance weighting (IDW) [63] method, because of its relatively fast and easy computation and interpretation [64]. IDW sums the values of nearby points multiplied by a weighting factor. The weights are a decreasing function of distance. For this purpose, the Arc GIS 10.1 software was used to make spatial distribution maps.

3. Results and Discussion

3.1. Annual Departure from Normal. The time series of total annual precipitation averaged over whole Beijing-Tianjin-Hebei metropolitan areas and the corresponding percent of normal precipitation index computed with respect to 1961–1990 climate normal are presented in Figure 2. The results showed that the year 1965 was the driest year during the observed period with 70.8% of normal precipitation. All of the stations had the precipitation below the annual mean precipitation, while the annual mean precipitation was 371.56 mm. The average precipitation for the observed period was 519.19 mm. In addition, it is necessary to emphasize another two years, which were severe and extremely dry across most of the area, compared to 1961–1990 climate normal: (1) 1965 with 70.8% of normal precipitation and (2) 1997 with 71.57% of normal precipitation.

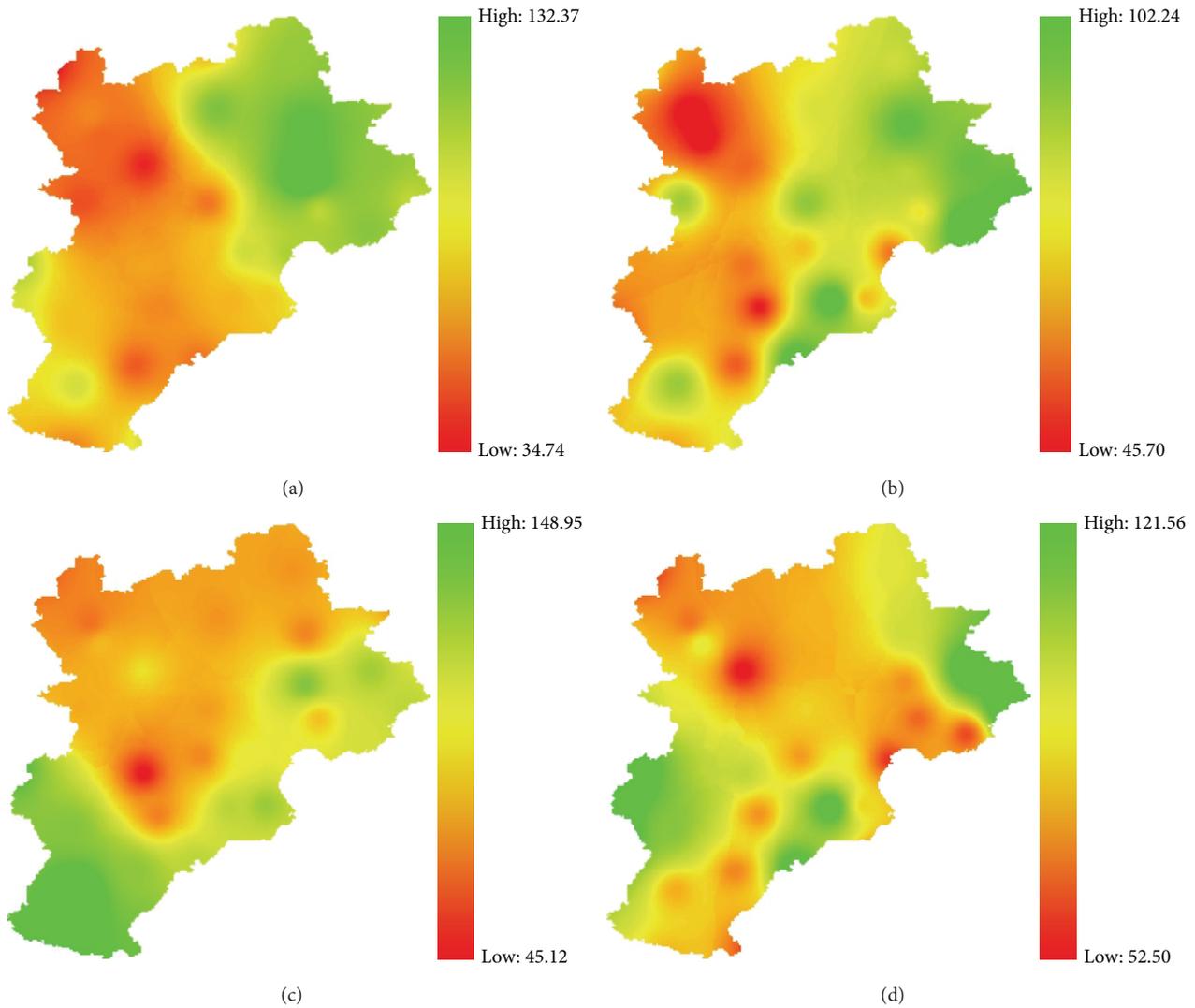


FIGURE 3: Spatial distribution of percentage of annual mean precipitation relative to 1961–1990 climate normal over the course of the years (a) 1965; (b) 1997; (c) 2000; and (d) 2002.

Further analysis showed that the year of 2000 had the summer that was extremely warm and extremely dry. And the summer of the year of 2000 was the warmest one during the last 50 years since the meteorological measurements have been conducted in BTHMA [65].

Spatial distribution of percentage of annual mean precipitation relative to 1961–1990 climate normal over the course of the three driest years (1965, 1997, and 2002) and the 2000 year is shown in Figure 3.

The average value at the region level was 68.54% during the 1965 year and 69.25% during the year 1997. In the course of the year 2002, most of the area had the percent of normal precipitation between 52.50% and 121.56%. At the same time, the RDI value of the whole area in 2002 was -0.95 , which was only above -1.00 for the 1997 year and -1.02 for the 1965 year. In addition, the year 2000 had the highest values, which varied between 45.12% and 148.95% with the highest average value 87.02%.

3.2. Analysis of Drought Indices

3.2.1. Analysis of RDI. Percentage of years affected by various drought severity levels in Beijing–Tianjin–Hebei metropolitan areas during the period 1960–2010 is illustrated in Figure 4(a). 66.91% of the frequency of drought belongs to the near normal drought category. The frequency of drought was 15.30%, while the frequency of wet periods was 17.79%. Percentage of Beijing–Tianjin–Hebei metropolitan areas affected by drought during the period 2000–2010 is presented in Figure 4(b). The frequency of drought was 20.35% during the period 2000–2010, which was above 15.30%. Therefore, the results reveal that drought condition during the years 2000 and 2010 was more severe than that during the years 1960–2010.

Spatial distribution of RDI-12 during the three driest years (1965, 1997, and 2002) and the 2000 year as the year with the warmest summer and the topmost PET is shown in Figure 5. The extreme drought occurred in some areas

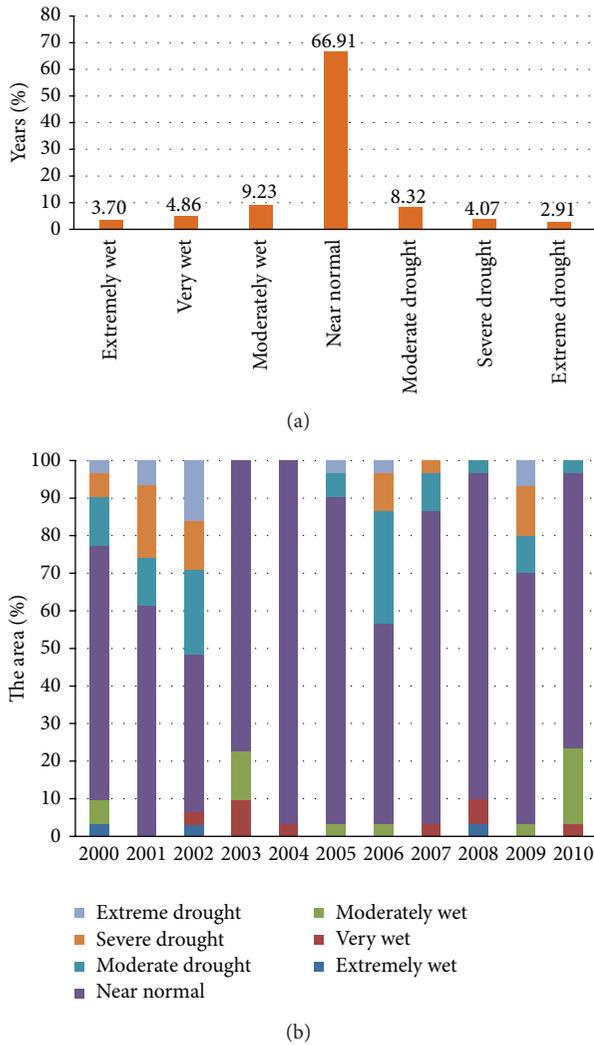


FIGURE 4: (a) Percentage of years affected by various drought severity levels, 1960–2010, and (b) percentage of BTHMA affected by drought, 2000–2010.

during the year 2000, 2001, 2002, 2005, 2006, and 2009, while drought of varying degrees occurred except the years 2003 and 2004. Furthermore, the majority of the area had severe and extreme droughts during the years 1965 and 1972.

3.3. **Drought Patterns.** The RDI-12 was used to identify drought patterns for the period 1960–2010. First, the Kaiser-Meyer-Olkin (KMO) and Bartlett’s tests were applied to these indices. The KMO measure of sampling adequacy was 0.786 for the RDI. High values of the KMO test (>0.50) suggest that the selected indices are adequate for the PCA. Bartlett’s test of sphericity has the P value < 0.0001 for $\alpha = 0.05$, which is good and it is an indication we can continue with the PCA. The first seven eigenvalues for the RDI-12 with the corresponding error bars at 5% significance level are shown in Figure 6. According to North’s rule of thumb and the scree plot of the eigenvalues, the first two principal components (PCs) were selected for varimax rotation to achieve more stable spatial

TABLE 3: Explained variance (%) by the loadings with and without rotation for the RDI-12 during 1960–2010.

Principal component	RDI-12	
	Unrotated (%)	Varimax rotated
PC-1	43.391	26.877
PC-2	10.080	26.594
Cumulative percentage of total variation	53.471	53.471

TABLE 4: Variance decomposition for the optimal classification.

	RDI-12	
	Absolute	Percent (%)
Between-classes	0.131	57.25
Within-class	0.067	42.75
Total	0.198	100.00

patterns. Table 3 summarizes the variances of unrotated and rotated components. The first unrotated component had the biggest variance value 43.391% for the RDI-12. The percentage of the cumulative variance for the RDI-12 was 53.471%. The results also show that the cumulative variance of the varimax rotated components remains unaltered with respect to the unrotated cases.

Scatter plot of the correlations between variables and PCs after varimax rotation for the RDI-12 is shown in Figure 7. Each variable is a point whose coordinates are given by the loadings on the PCs. The correlation for the RDI indices is positive. In particular, the station of Tianjin is located by the Bohai Sea, of which the time series of precipitation and evaporation were extremely different from the other stations. Therefore, the station of Tianjin was apart from the group in Figure 7.

In Figure 8 the spatial patterns of varimax rotated loadings (R-Loading 1 and R-Loading 2) of the RDI-12 for the period 1960–2010 are shown. Further, the time variability of the RPCs of the RDI-12 and the corresponding linear trend are presented too. The quite small linear trends are identified. The remarkable dry events of different magnitudes are identified during the following periods: 1965–1967, 1971–1975, 1978–1982, 1989–1993, 1997–2002, and 2005–2007. Both the RPC-1 and the RPC-2 showed that the worst drought event occurred in the year 1965.

The R-Loadings seem to localize well in space three distinct subregions, the northern, southern, and northeastern part of BTHMA. The identified subregions are characterized by different drought variability that depends on the different precipitation regimes in these areas.

The agglomerative hierarchical cluster analysis (AHCA) was used to investigate drought patterns by grouping observations into clusters. It was applied to the RPCs using Ward’s method and Euclidean distance. The goal was to search an optimal grouping for which the observed values within each cluster were similar, while the clusters were dissimilar to each other. The obtained variance decomposition for the optimal classification is summarized in Table 4. Since

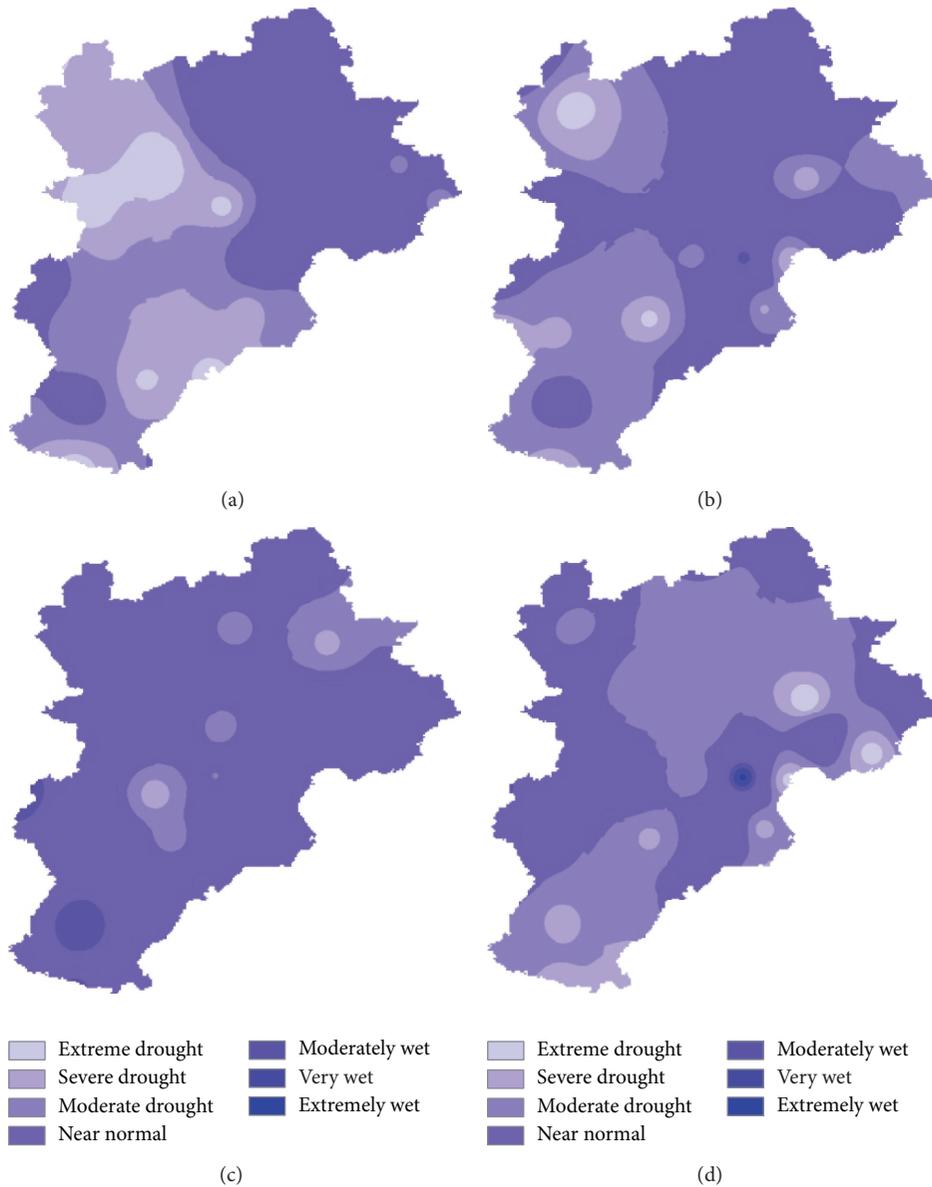


FIGURE 5: Spatial distribution of RDI-12 over the course of the years (a) 1965; (b) 1997; (c) 2000; and (d) 2002.

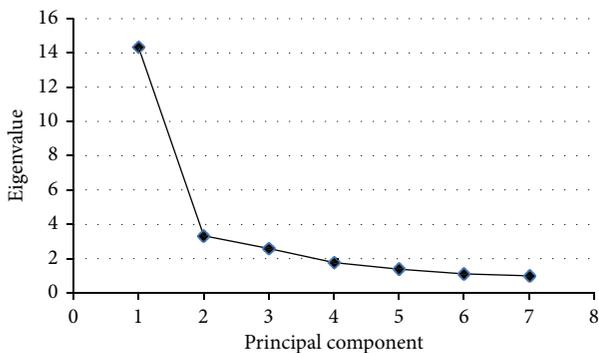


FIGURE 6: Scree plot for the first seven eigenvalues and the corresponding error bars at 5% significance level resulting from the S-mode PCA applied to (a) the RDI.

between-cluster variation is much larger than within cluster variation, thus, obtained PCs successfully reflect the cluster structure. Applying the AHCA, three distinct subregions (R1 with 3 stations; R2 with 12 stations; and R3 with 18 stations) were identified (Figure 9).

3.3.1. Drought Characteristics of the Identified Subregions.

Detected subregions are defined as follows: (1) region R1 is located in southeast part of BTHMA; (2) region R2 is located in the north of BTHMA; and (3) region R3 is located in the southern part of BTHMA.

The R1 is mostly forested with the average annual precipitation to 800 mm, while the R2 is characterized by a moderate-precipitation regime with the average annual precipitation to 650 mm. The R3 is characterized by the lowest

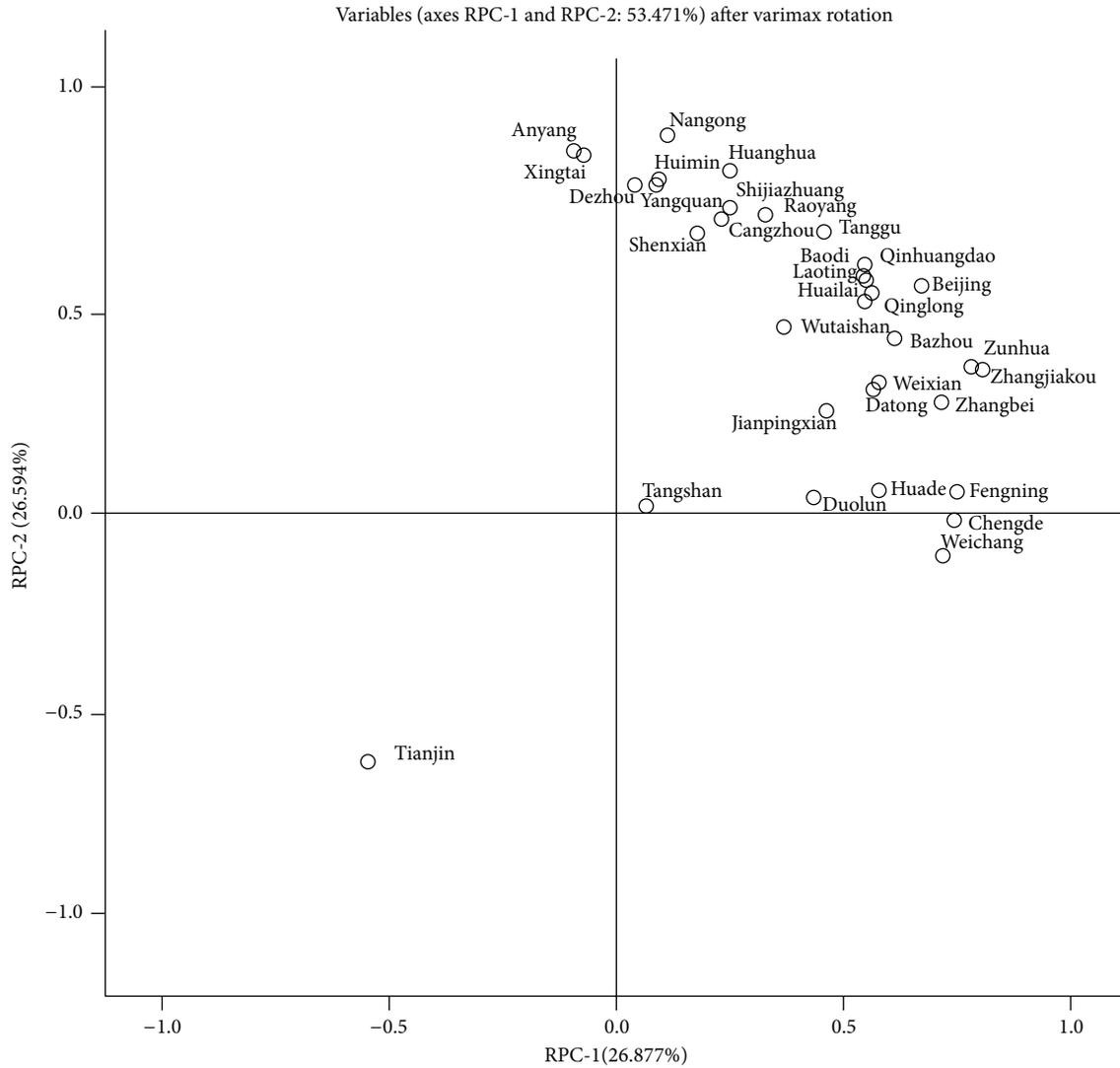


FIGURE 7: Scatter plot of the correlations between variables and principal components after varimax rotation for the RDI-12 series.

amount of precipitation in the area and mostly intensive agriculture. The R1 and R2 had the monthly precipitation values above average, while R3 had the precipitation values under average of BTHMA.

The diversity of time variability of the regional RDI-12 for the three subregions shown in Figure 10 was existed. Among them, the value of regional RDI-12 for R2 was above that for R3 and below that for R1.

In this respect, based on the RDI-12 values and defined categories of dry and wet conditions (Table 2), the periods of drought were 1965–1967, 1971–1975, 1978–1982, 1989–1993, 1997–2002, and 2005–2007, whereas the periods with wet conditions were 1960–1964, 1968–1970, 1966–1977, 1983–1988, 1994–1996, 2003–2004, and 2008–2010. The distributions of the dry and wet years are as follows: R1-29 (4) dry (wet) years, R2-25 (8) dry (wet) years, and R3-21 (11) dry (wet) years. The year 1965 was the driest, while 1964 was the wettest during the observed period. The evolution of drought in region R2 is similar to region R1.

The obtained results are in line with conclusions reported by Yuan and Wu (2000) [28], Yan et al. (2010) [29], and Liu et al. (2012) [8].

3.3.2. *Frequency of Drought.* The drought occurrence is analyzed based on Table 2 that defines drought classes related to the RDI values. In this respect, the frequency distribution of the RDI-12 values was divided in seven classes. The ratio between the number of drought occurrences in each of RDI classes and the total number of events counted for all stations in a given region was represented as the frequency of drought. The percentage of drought and wet occurrence expressed in seven classes of RDI-12 drought categories for each individual region for the period 1960–2010 is illustrated in Figure 11. It should be noted that the frequency distribution of the RDI-12 values expressed in percent was closely similar for the three identified regions.

Spatial distribution of the frequency of the RDI-12 values expressed in percent related to the moderate, severe, and

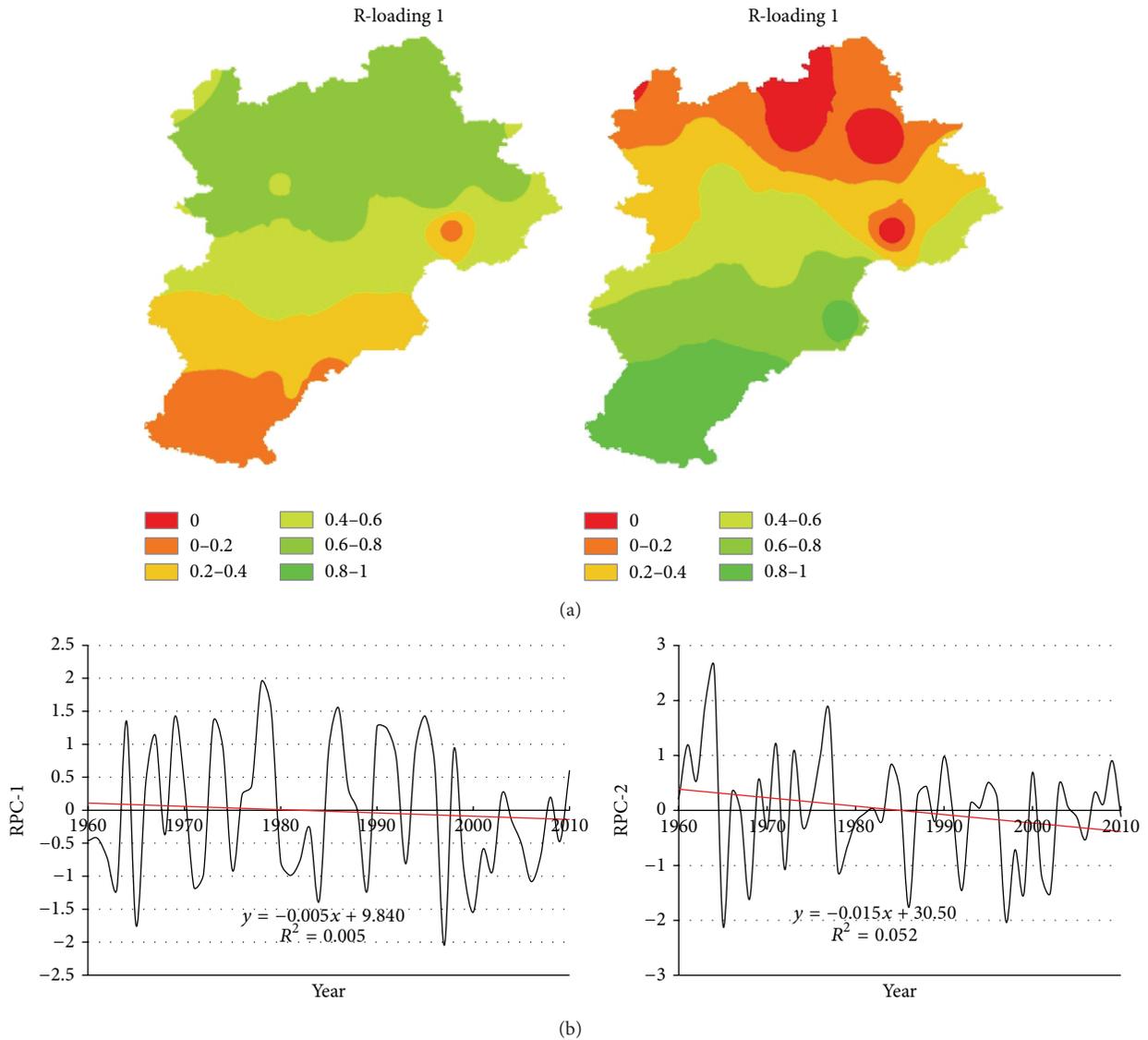


FIGURE 8: (a) Varimax rotated loading patterns (R-loading) of the RDI-12 and (b) time variability of the rotated PC scores (RPC) of the RDI-12 for the period 1960–2010 and the corresponding linear trend.

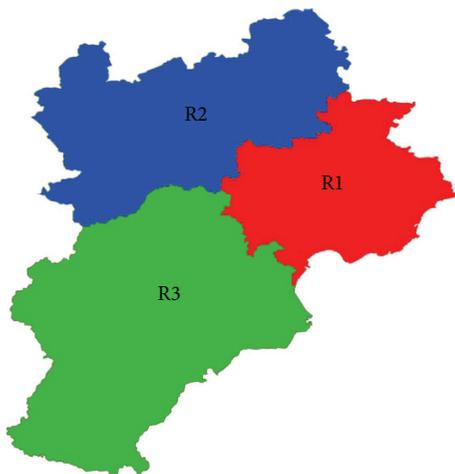


FIGURE 9: Drought-based regionalization in Beijing-Tianjin-Hebei metropolitan areas.

extreme drought is shown in Figure 12. The distribution is based on the station values of the RDI-12 values. The frequency of moderate drought ranged between 1.96% and 17.65%. The majority of the area had the frequency between 7% and 9%. The highest frequency of severe drought occurrence was 7.84%, which was evenly located in the regions R2 and R3. The average at the area level was 4.69%. The frequency of severe drought ranged between 0% and 7.84%. According to the spatial distribution of the frequencies of the RDI-12 values, the highest frequency of extreme drought (7.84%) was detected at Nangong station (in region R3). Region R1 had the frequency of extreme drought ranged between 0% and 2.94%.

4. Conclusions

The drought was investigated in Beijing-Tianjin-Hebei metropolitan areas using monthly precipitation time series

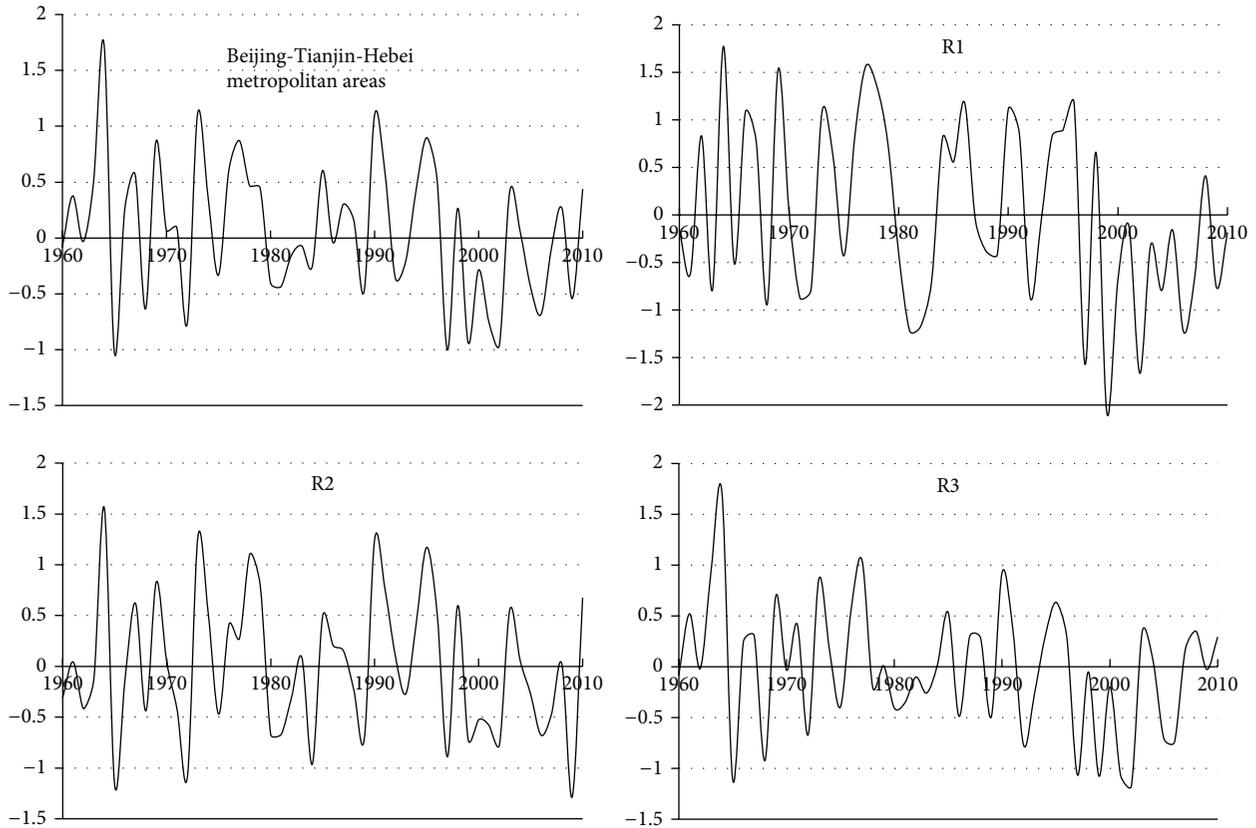


FIGURE 10: Time series of RDI-12 for BTHMA and three subregions.

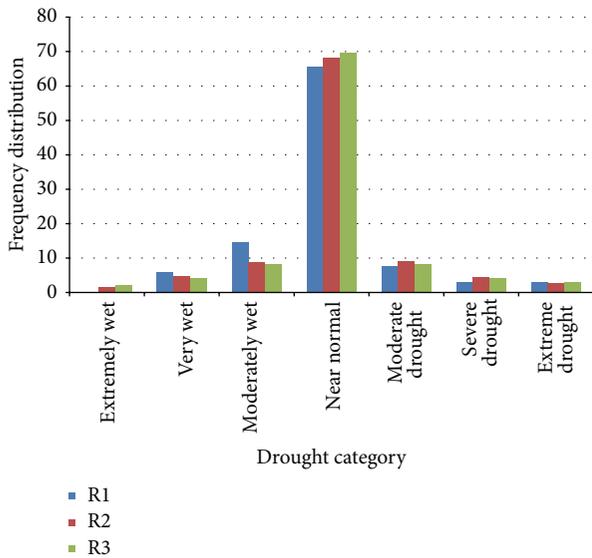


FIGURE 11: Frequency distribution of the RDI-12 values for each individual region during the period 1960–2010.

from 33 stations during the period of 1960–2010. The temporal and spatial patterns of drought were analyzed by applying the S-mode PCA to the RDI estimated on 12-month timescales. According to the error bars of the North rule of thumb and scree plot, two principal components were

retained. These components were localized well in space three distinct subregions, characterizing by different drought variability. The AHCA confirmed the results from PCA analysis and identified three different drought subregions R1, R2, and R3, which are located in southeast, north, and south of BTHMA, respectively. The results of both the PCA and the AHCA analysis obtained a very similar time variability of the regional RDI-12.

The characteristics of drought were analyzed in terms of the temporal evolution of the RDI-12 values and the frequency of drought at the area level and for three regions. The linear regression method was used for time variability analysis of drought in each identified subregion as well as for the whole area. The frequency of drought was 9.10%, while the distribution of wet periods was 27.27% in the given regions. 66.91% of the frequency of drought belongs to the near normal drought category. According to the RDI-12, the average number of the dry years in the detected regions was about 8 years during the period 1960–2010. The year 1965 was the driest, while 1964 was the wettest during the observed period. Three years (1965, 1997, and 2002) were detected as the severe and extremely dry in the majority of the area and analyzed by the percent of normal precipitation index computed with respect to 1960–1990 climate normal.

The obtained results provide support information to improve water resources management in the study area. Further research should be performed to detect the trends

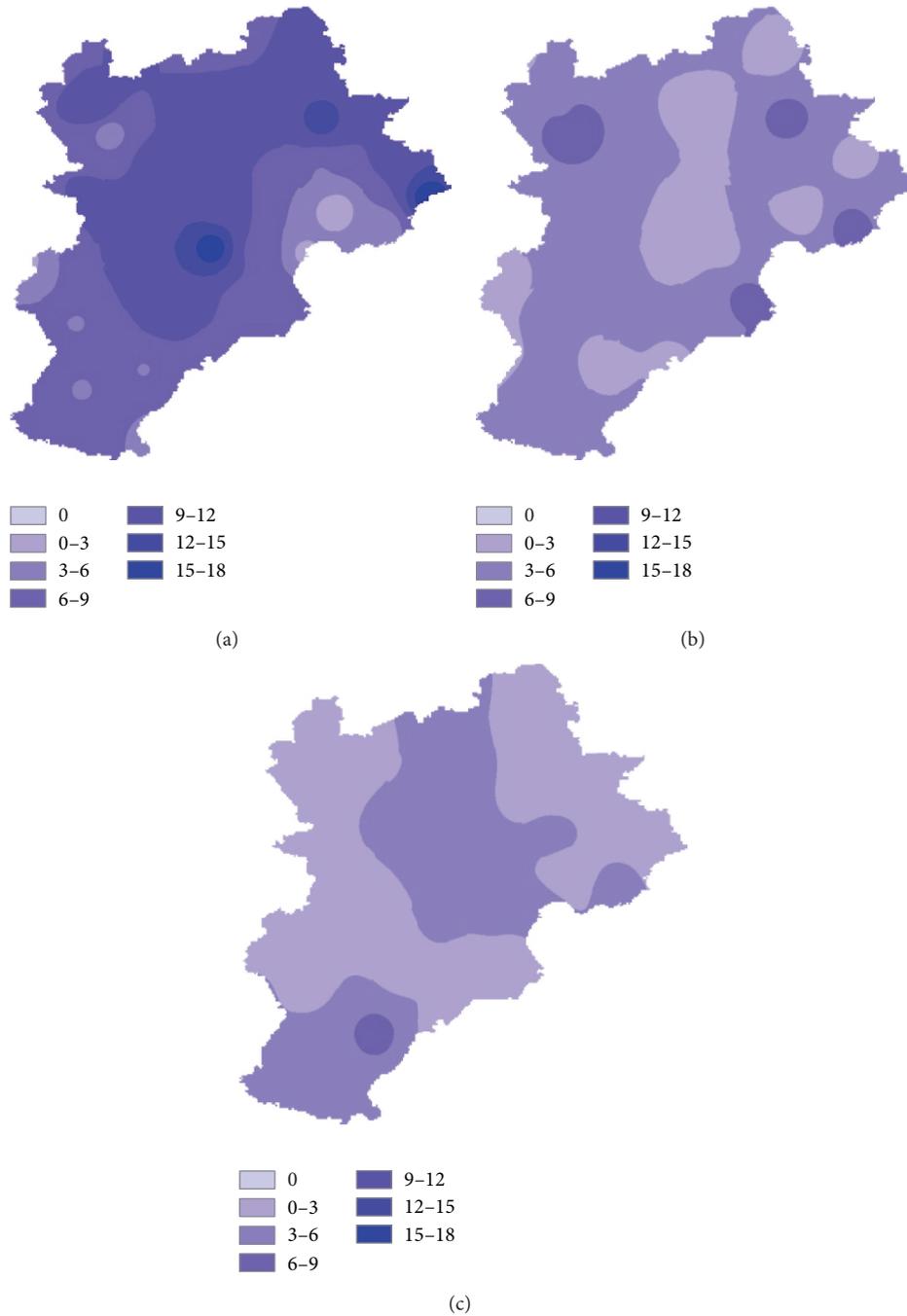


FIGURE 12: Spatial distribution of the frequency (in percent) of the RDI-12 values related to the following categories: (a) moderate, (b) severe, and (c) extreme drought.

of drought in Beijing-Tianjin-Hebei metropolitan areas and comparative analysis of the drought indices based on precipitation and evapotranspiration and their impact on agricultural production.

Conflict of Interests

The authors declare that there is no conflict of interests regarding the publication of this paper.

Acknowledgments

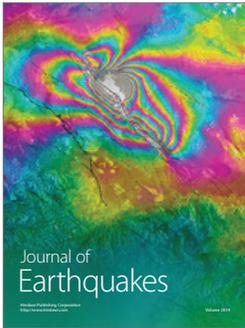
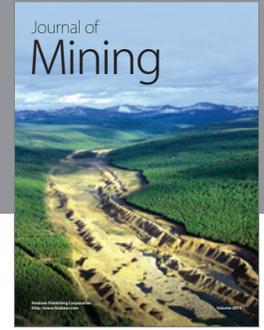
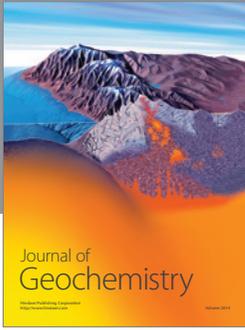
The research was financed by the National Science and Technology Support Plan during the 12th Five-year Plan Period of China (No. 2012BAC19B03 and 2013BAC10B01). It is also realized as a part of the Project “Statistical Analysis and Information Extraction of High-dimensional Complicate Data” jointly funded by Scientific Research Project of Beijing Educational Committee (No. KZ201410028030). The authors

would like to thank the Editors and the anonymous reviewers for their crucial comments, which improved the quality of this paper.

References

- [1] J. S. Wang, J. Y. Guo, Y. W. Zhou, and L. F. Yang, "Progress and prospect on drought indices research," *Arid Land Geography*, vol. 30, pp. 60–65, 2007.
- [2] Y. B. Yao, A. X. Dong, Y. R. Wang, X. Y. Zhang, and J. H. Yang, "Compare research of the regional arid characteristic base on Palmer drought severity index in spring over China," *Arid Land Geography*, vol. 30, pp. 22–29, 2007.
- [3] W. Yu, M. Shao, M. Ren, H. Zhou, Z. Jiang, and D. Li, "Analysis on spatial and temporal characteristics drought of Yunnan Province," *Acta Ecologica Sinica*, vol. 33, pp. 317–324, 2013.
- [4] J. Yoo, H.-H. Kwon, T.-W. Kim, and J.-H. Ahn, "Drought frequency analysis using cluster analysis and bivariate probability distribution," *Journal of Hydrology*, vol. 420–421, pp. 102–111, 2012.
- [5] T. Sternberg, "Regional drought has a global impact," *Nature*, vol. 472, no. 7342, p. 169, 2011.
- [6] A. A. Paulo, R. D. Rosa, and L. S. Pereira, "Climate trends and behaviour of drought indices based on precipitation and evapotranspiration in Portugal," *Natural Hazards and Earth System Science*, vol. 12, no. 5, pp. 1481–1491, 2012.
- [7] M. Gocic and S. Trajkovic, "Spatiotemporal characteristics of drought in Serbia," *Journal of Hydrology*, vol. 510, pp. 110–123, 2014.
- [8] X. Y. Liu, D. L. Li, and J. S. Wang, "Spatiotemporal characteristics of drought over China during 1961–2009," *Journal of Desert Research*, vol. 32, pp. 473–483, 2012.
- [9] S. Y. Li, R. H. Liu, L. K. Shi, and Z. H. Ma, "Analysis on drought characteristic of Henan in recent 40 years based on meteorological drought composite index," *Journal of Arid Meteorology*, vol. 27, pp. 97–102, 2009.
- [10] Y. Bao, C. Meng, S. Shen, X. Qiu, P. Gao, and C. Liu, "Temporal and spatial patterns of droughts for recent 50 years in Jiangsu based on meteorological drought composite index," *Acta Geographica Sinica*, vol. 66, no. 5, pp. 599–608, 2011.
- [11] A. K. Mishra and V. P. Singh, "A review of drought concepts," *Journal of Hydrology*, vol. 391, no. 1–2, pp. 202–216, 2010.
- [12] A. Dai, "Characteristics and trends in various forms of the palmer drought severity index during 1900–2008," *Journal of Geophysical Research D: Atmospheres*, vol. 116, no. 12, Article ID D12115, 2011.
- [13] L. Vasiliades, A. Loukas, and N. Liberis, "A water balance derived drought index for Pinios river basin, Greece," *Water Resources Management*, vol. 25, no. 4, pp. 1087–1101, 2011.
- [14] T. Gebrehiwot, A. van der Veen, and B. Maathuis, "Spatial and temporal assessment of drought in the Northern highlands of Ethiopia," *International Journal of Applied Earth Observation and Geoinformation*, vol. 13, no. 3, pp. 309–321, 2011.
- [15] H. Vangelis, D. Tigkas, and G. Tsakiris, "The effect of PET method on Reconnaissance Drought Index (RDI) calculation," *Journal of Arid Environments*, vol. 88, pp. 130–140, 2013.
- [16] R. R. Heim Jr., "A review of twentieth-century drought indices used in the United States," *Bulletin of the American Meteorological Society*, vol. 83, no. 8, pp. 1149–1165, 2002.
- [17] G. Tsakiris and H. Vangelis, "Establishing a drought index incorporating evapotranspiration," *European Water*, vol. 9, pp. 3–11, 2005.
- [18] G. Tsakiris, D. Pangalou, and H. Vangelis, "Regional drought assessment based on the reconnaissance drought index (RDI)," *Water Resources Management*, vol. 21, no. 5, pp. 821–833, 2007.
- [19] M. A. A. Zarch, H. Malekinezhad, M. H. Mobin, M. T. Dastorani, and M. R. Kousari, "Drought monitoring by reconnaissance drought index (RDI) in Iran," *Water Resources Management*, vol. 25, no. 13, pp. 3485–3504, 2011.
- [20] N. A. Elagib and M. M. Elhag, "Major climate indicators of ongoing drought in Sudan," *Journal of Hydrology*, vol. 409, no. 3–4, pp. 612–625, 2011.
- [21] N. Farajalla and R. Ziade, "Drought frequency under a changing climate in the Eastern Mediterranean: the Bekaa Valley, Lebanon," *Geophysical Research Abstracts*, vol. 12, p. 116, 2010.
- [22] D. Khalili, T. Farnoud, H. Jamshidi, A. A. Kamgar-Haghighi, and S. Zand-Parsa, "Comparability analyses of the SPI and RDI meteorological drought indices in different climatic zones," *Water Resources Management*, vol. 25, no. 6, pp. 1737–1757, 2011.
- [23] D. G. C. Kirono, D. M. Kent, K. J. Hennessy, and F. Mpelasoka, "Characteristics of Australian droughts under enhanced greenhouse conditions: results from 14 global climate models," *Journal of Arid Environments*, vol. 75, no. 6, pp. 566–575, 2011.
- [24] G. Rossi, M. Benedini, G. Tsakiris, and S. Giakoumakis, "On regional drought estimation and analysis," *Water Resources Management*, vol. 6, no. 4, pp. 249–277, 1992.
- [25] B. Clausen and C. P. Pearson, "Regional frequency analysis of annual maximum streamflow drought," *Journal of Hydrology*, vol. 173, no. 1–4, pp. 111–130, 1995.
- [26] T. Estrela, M. Menendez, M. Dimas et al., "Sustainable water use in Europe. Part 3: extreme hydrological events: floods and droughts," Environmental issue report no. 21, vol. 84, European Environment Agency, Copenhagen, Denmark, 2001.
- [27] C. Yoo and S. Kim, "EOF analysis of surface soil moisture field variability," *Advances in Water Resources*, vol. 27, no. 8, pp. 831–842, 2004.
- [28] G. F. Yuan and L. H. Wu, "Spatial and temporal analysis on agricultural drought under double cropping system in Beijing-Tianjin-Hebei plain area," *Chinese Journal of Agrometeorology*, vol. 21, pp. 5–9, 2000.
- [29] F. Yan, Y. J. Wang, and B. Wu, "Spatial and temporal distributions of drought in Hebei Province over the past 50 years," *Geographical Research*, vol. 39, pp. 423–430, 2010.
- [30] X. Y. Liu, D. L. Li, and J. S. Wang, "Spatiotemporal characteristics of drought over China during 1961–2009," *Journal of desert research*, vol. 32, pp. 473–483, 2012.
- [31] G. Willeke, J. R. M. Hosking, J. R. Wallis, and N. B. Guttman, "The National Drought Atlas," Institute for Water Resources Report 94-NDS-4, Army Corps of Engineers, U.S., 1994.
- [32] G. Tsakiris, "Meteorological drought assessment," in *Proceedings of the Needs of the European Research Program MEDROPLAN (Mediterranean Drought Preparedness and Mitigation Planning)*, Zaragoza, Spain, 2004.
- [33] G. Tsakiris and H. Vangelis, "Establishing a drought index incorporating evapotranspiration," *European Water*, vol. 9–10, pp. 3–11, 2005.
- [34] P. Droogers and R. G. Allen, "Estimating reference evapotranspiration under inaccurate data conditions," *Irrigation and Drainage Systems*, vol. 16, no. 1, pp. 33–45, 2002.

- [35] H. Vangelis, M. Spiliotis, and G. Tsakiris, "Drought severity assessment based on bivariate probability analysis," *Water Resources Management*, vol. 25, no. 1, pp. 357–371, 2010.
- [36] G. L. Hargreaves and Z. A. Samani, "Reference crop evapotranspiration from temperature," *Applied Engineering in Agriculture*, vol. 1, pp. 96–99, 1985.
- [37] C. W. Thornthwaite, "An approach toward a rational classification of climate," *Geographical Review*, vol. 38, pp. 55–94, 1948.
- [38] A. R. Pereira and W. O. Adaptation, "Adaptation of the Thornthwaite scheme for estimating daily reference evapotranspiration," *Agricultural Water Management*, vol. 66, no. 3, pp. 251–257, 2004.
- [39] J. Luo and Y. Rong, "Evaluation on several Empirical methods for estimating potential evapotranspiration in North China," 2006, <http://www.paper.edu.cn>.
- [40] M. Abramowitz and I. A. Stegun, *Handbook of Mathematical Functions*, Dover, New York, NY, USA, 1965.
- [41] X. Shao, F. Yao, and J. Zhang, "Standardized precipitation evapotranspiration index shows drought research in North China," *Meteorological Monthly*, vol. 39, pp. 1154–1162, 2013.
- [42] D. Tigkas, H. Vangelis, and G. Tsakiris, "Drought and climatic change impact on streamflow in small watersheds," *Science of the Total Environment*, vol. 440, pp. 33–41, 2012.
- [43] H. Hotelling, "Analysis of a complex of statistical variables into principal components," *Journal of Educational Psychology*, vol. 24, no. 6, pp. 417–441, 1933.
- [44] M. R. Anderberg, *Cluster Analysis for Applications*, Academic Press, New York, NY, USA, 1973.
- [45] B. S. Everitt, *Cluster Analysis*, Edward Arnold, London, UK, 3rd edition, 1993.
- [46] L. S. Kalkstein and J. A. Skindlov, "An evaluation of three clustering procedures for use in synoptic climatological classification," *Journal of Climate & Applied Meteorology*, vol. 26, no. 6, pp. 717–730, 1987.
- [47] R. G. Fovell and M.-Y. C. Fovell, "Climate zones of the conterminous United States defined using cluster analysis," *Journal of Climate*, vol. 6, no. 11, pp. 2103–2135, 1993.
- [48] M. Widmann and C. Schar, "A principal component and long-term trend analysis of daily precipitation in Switzerland," *International Journal of Climatology*, vol. 17, no. 12, pp. 1333–1356, 1997.
- [49] Y. Dinpashoh, A. Fakhri-Fard, M. Moghaddam, S. Jahanbakhsh, and M. Mirnia, "Selection of variables for the purpose of regionalization of Iran's precipitation climate using multivariate methods," *Journal of Hydrology*, vol. 297, no. 1–4, pp. 109–123, 2004.
- [50] T. Raziei, B. Saghafian, A. A. Paulo, L. S. Pereira, and I. Bordi, "Spatial patterns and temporal variability of drought in Western Iran," *Water Resources Management*, vol. 23, no. 3, pp. 439–455, 2009.
- [51] D. S. Martins, T. Raziei, A. A. Paulo, and L. S. Pereira, "Spatial and temporal variability of precipitation and drought in Portugal," *Natural Hazards and Earth System Science*, vol. 12, no. 5, pp. 1493–1501, 2012.
- [52] I. T. Jolliffe, *Principal Component Analysis*, Springer, New York, NY, USA, 2002.
- [53] H. Abdi and L. J. Williams, "Principal component analysis," *Wiley Interdisciplinary Reviews: Computational Statistics*, vol. 2, no. 4, pp. 433–459, 2010.
- [54] U. Demšar, P. Harris, C. Brunson, A. S. Fotheringham, and S. McLoone, "Principal component analysis on spatial data: an overview," *Annals of the Association of American Geographers*, vol. 103, no. 1, pp. 106–128, 2013.
- [55] M. B. Richman, "Rotation of principal components," *Journal of Climatology*, vol. 6, no. 3, pp. 293–335, 1986.
- [56] G. R. North, T. L. Bell, R. F. Cahalan, and F. J. Moeng, "Sampling errors in the estimation of empirical orthogonal functions," *Monthly Weather Review*, vol. 110, pp. 699–706, 1982.
- [57] R. B. Cattell, "The scree test for the number of factors," *Multivariate Behavioral Research*, vol. 1, pp. 245–276, 1966.
- [58] M. S. Bartlett, "A note on the multiplying factors for various χ^2 approximations," *Journal of the Royal Statistical Society B: Methodological*, vol. 16, no. 2, pp. 296–298, 1954.
- [59] H. F. Kaiser, "A second generation little jiffy," *Psychometrika*, vol. 35, no. 4, pp. 401–415, 1970.
- [60] J. Ward, "Hierarchical grouping to optimize an objective function," *The Journal of the American Statistical Association*, vol. 58, pp. 236–244, 1963.
- [61] T. Raziei, I. Bordi, and L. S. Pereira, "A precipitation-based regionalization for Western Iran and regional drought variability," *Hydrology and Earth System Sciences*, vol. 12, no. 6, pp. 1309–1321, 2008.
- [62] J. F. Santos, I. Pulido-Calvo, and M. M. Portela, "Spatial and temporal variability of droughts in Portugal," *Water Resources Research*, vol. 46, no. 3, Article ID W03503, 2010.
- [63] D. Shepard, "A two-dimensional interpolation function for irregularly-spaced data," in *Proceedings of the ACM 23rd National Conference (ACM '68)*, pp. 517–524, 1968.
- [64] G. Y. Lu and D. W. Wong, "An adaptive inverse-distance weighting spatial interpolation technique," *Computers & Geosciences*, vol. 34, no. 9, pp. 1044–1055, 2008.
- [65] H. B. Shi, "Analysis of trends in the variability of summer high temperature days during 1960–2008 in Beijing, Tianjin and Hebei," *Meteorological Monthly*, vol. 37, pp. 1277–1282, 2011.



Hindawi

Submit your manuscripts at
<http://www.hindawi.com>

